Lecture 10 – Carbonate chemistry and acid base

(1) Many of the Lakes in New England have become acidified due to industrial emissions.

The pH of one such lake is 5.2. In descending order of concentration (highest to lowest), the inorganic carbon speciation in this lake would be:

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a. H<sub>2</sub>CO<sub>3</sub>, HCO<sup>3-</sup>, CO<sub>3</sub><sup>2-</sup>
b. HCO<sup>3-</sup>, CO<sub>3</sub><sup>2-</sup>, H<sub>2</sub>CO<sub>3</sub>
c. CO<sub>3</sub><sup>2-</sup>, HCO<sup>3-</sup>, H2CO<sub>3</sub>
d. HCO<sub>3</sub>-, H<sub>2</sub>CO<sub>3</sub>, CO<sub>3</sub><sup>2-</sup>
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The ocean pH is regulated by:

- (a) the conversion of sulfate to sulfide
- (b) the exchange of nitrogen with the atmosphere
- (c) the solution and precipitation of silicate minerals in seawater
- (d) the carbonate system
- (e) non of these

Suppose you bubble a seawater sample with N₂ gas until a white precipitate forms.

Which of the following statements is most likely **true**?

- 1. Alkalinity is unchanged
- 2. The precipitate is sodium bicarbonate (NaHCO₃)
- 3. Total CO_2 (ΣCO_2) is unchanged
- 4. The precipitate is calcium carbonate (CaCO₃)
- 5. pH goes down

As a result of CaCO₃ precipitation:

- a. Σ_{CO2} goes down $\text{Ca}^{2+} + \text{CO}_3^{2-} = \text{CaCO}_3$
- b. pH goes up
- c. P_{CO2} goes down
- d. Alkalinity stays constant
- e. respiration goes up

When (Alk- Σ CO₂) is high

- (a) pH is low
- (b) pH is high more CO_3^2 relative to HCO_3^- ($HCO_3^-+2CO_3^2$) ($HCO_3^-+CO_3^2$)
- (c) pH is not related to this value
- (d) more CaCO3 will dissolve
- (e) photosynthesis in low

- (2) What is the alkalinity of each of the following solutions?
- a) One mole of MgSO₄ dissolved in 400 liters of CO₂-free distilled water.
- b) One mole of Na₂CO₃ dissolved in 1000 liters of CO₂-free distilled water.
- c) One mole of KCl dissolved in 5000 liters of CO₂-free distilled water.
- d) One mole of NaCl and 2 moles of Ca(HCO₃)₂ in 1000 liters of CO₂-free distilled water.
- e) Average deep ocean water, where $[CO_3^{2-}] = 90 \mu M$, $[HCO_3^{-}] = 2350 \mu M$, $H_2CO_3 + CO_2 = 45 \mu M$. You may calculate the carbonate alkalinity and neglect other species.
- a) and c) ALK = 0; no change in carbonate bicarbonate or borate
- b) There is one mol of CO_3^2 in 1000 L, which is 2 equivalents in 1000 liters or 2 meg/l
- d) 2x2 = 4 equivalents HCO_3^- in 1000 liters = 4 meg/l
- e) Alkalinity = $HCO_3^- + 2CO_3^{2-} + OH^- H^+$

We can usually approximate alkalinity by neglecting OH^- and H^+ since they are relatively small, and almost equal to each other in seawater. So $2350 + 2 \times 90 = 2440 \text{ meq/l}$ Note that CO_2 and H_2CO_3 do not affect alkalinity and never will.

- (3) The deep water of the Black Sea does not get replenished very often thus dissolved oxygen is absent and hydrogen sulfide is very high.
 - a. What is the sulfide (S^{2-}) concentration in the deep water of the Black Sea. Use the graphical or algebraic approach (your choice).

Total sulfide (ST) equals $400\mu M$; pH is 7.7 There are two apparent acidity constants for H_2S in seawater: $K_1' = 10^{-7}$ and $K_2' = 10^{-14}$

The equations to consider:

$$H_2S = H^+ + HS^ K_1' = 10^{-7}$$
 $K_1 = [H^+] [HS^-] / [H_2S]$ $[HS^-] = K_1 [H_2S]/a_{H^+}$
 $HS^- = H^+ + S^{2-}$ $K_2' = 10^{-14}$ $K_2 = [H^+] [S^{2-}] / [HS^-]$ $[S^-] = K_2 [HS^-]/a_{H^+}$
 $Total S = H_2S + HS^- + S^{2-} = 400 \times 10^{-6} M = 10^{-3.39}$

For this we need to first calculate the S^{2-} concentration at these conditions $S_T = [S^{2-}] \{a_{H+}^2/K_1K_2 + a_{H+}/K_2 + 1\}$ $[S^{2-}] = S_T/\{a_{H+}^2/K_1K_2 + a_{H+}/K_2 + 1\} = 1.67 \times 10^{-10}$

b. Is iron sulfide (FeS) supersaturated? Remember that the solubility reaction is: $FeS = Fe^{2+} + S^{2-}$

The apparent solubility constant in seawater is: $K'sp = 10^{-19}$ The concentration of Fe^{2+} is 200 nM

 $200 \times 10^{-9} \times 1.67 \times 10^{-10} = 3.34 \times 10^{-17} > 10^{-19}$ so the water is supersaturated with respect to FeS

- (4) I made a 1liter solution with 1 mole NaHCO₃ and 0.5 moles Na₂CO₃.
 - (a) What is the alkalinity of this solution?
 - (b) If no CO_2 escaped what is the DIC?
 - (c) What are the two major carbonate species in solution?

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(a) Alk = [HCO_3^-] + 2[CO_3^2] 1 mole HCO_3^{-3} and \frac{1}{2} mole CO_3^{-2} so it is 1 + 2 \times \frac{1}{2} = 2 equivalents
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- (b) $C_T = DIC = [HCO_3^-] + [CO3^2] = 1 + 0.5 = 1.5 \text{ moles/liter}$
- (c) $[HCO_3]$ and $[CO_3^2]$

a sink for CO_2 (on short time scales).

(5) The Alk of a sample of surface water is 2.35 meq/l and its ΣCO_2 is 2.15 mmol/l. The Alk of a sample from 4 km depth at the same location are 2.45meq/l and the ΣCO_2 is 2.40 mmol/l. What are the pH and $[CO_3^{2-}]$ for each sample? What is the saturation index for these samples with respect to calcite? Ksp =4.5 x 10^{-7} , (Ca) = 1.03 x 10^{-2} M

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Alk = [HCO<sub>3</sub><sup>-</sup>] + 2[CO<sub>3</sub><sup>2</sup>-]; DIC = [HCO<sub>3</sub><sup>-</sup>] + [CO<sub>3</sub><sup>2</sup>-];

[HCO<sub>3</sub><sup>-</sup>] = 2[DIC] - Alk [CO<sub>3</sub><sup>2</sup>-] = Alk-DIC

[HCO<sub>3</sub><sup>-</sup>] = [H<sup>+</sup>] + [CO<sub>3</sub><sup>2</sup>-] K<sub>2</sub> = 10<sup>-9</sup> K<sub>2</sub>= [H<sup>+</sup>] x [CO<sub>3</sub><sup>2</sup>-] / [HCO<sub>3</sub><sup>-</sup>] H<sup>+</sup> = K<sub>2</sub> [HCO<sub>3</sub><sup>-</sup>]/[CO<sub>3</sub><sup>2</sup>-]
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So: For surface water: $[CO_3^{2-}] = 200 \,\mu\text{M}$; $[HCO_3^{-}] = 1.95 \,\text{mM}$; and pH = 8.0 For deep water: $[CO_3^{2-}] = 50 \,\mu\text{M}$; $[HCO_3^{-}] = 2.35 \,\text{mM}$; and pH = 7.3

$$(\text{Ca}^{2+})$$
 (CO_3^{2-}) = surface: 1.03 x 10⁻² x 200 x 10⁻⁶ = 2.06 x 10⁻⁶ > 4.5 x 10⁻⁷ super saturated (Ca^{2+}) (CO_3^{2-}) = deep: 1.03 x 10⁻² x 50 x 10⁻⁶ = 5.15 x 10⁻⁷ > 4.5 x 10⁻⁷ super saturated but not by much

- (6) The popular press and a few scientists claim that the health of coral reefs is crucial to humanity because reef CaCO₃ represents a sink for anthropogenic CO₂.
- a. Explain the qualitative effect of $CaCO_3$ formation in reefs (skeletogenesis) on alkalinity and ΣCO_2

 $Ca^{2+} + CO_3^{2-} == CaCO_3$ $Ca^{2+} + 2HCO_3^{-} == CaCO_3 + CO_2 + H_2O$ CaCO₃ ppt. Alk goes down, DIC goes down, Alk change = 2xDIC change; But the on a short term this could actually be a source of atmospheric CO_2 since HCO_3^{-} is used and not atmospheric $CO_2!!$

b. The expression for P_{CO2} in terms of alkalinity and ΣCO_2 is: $P_{CO2} = \frac{(2\Sigma CO_2 - Alk)^2}{K'(Alk - \Sigma CO_2)}$

Use this expression to support or refute this idea (e.g. that reefs are a sink for CO_2) You may find it useful to plug demonstration numbers into this equation for comparison purposes. A typical surface value for ΣCO_2 is 2.0 mM and for Alk is 2.2 meq/l. What value of P_{CO2} is this water in equilibrium with and how would it change if about 0.1 mM of ΣCO_2 was removed to form reef material. $K' = 4.5 \times 10^4$ mmol I^{-1} atm⁻¹. Using the above equation and keeping in mind that 0.1 mM change in DIC will correspond to 0.2 meq/l change in Alk. You will find that 0.1 mM reduction in DIC thus 0.2 meq/l reduction on Alk will result in higher P_{CO2} thus this does not support the reef as

(7) A sample of seawater has a carbonate alkalinity of 2400 μ eq/l and TCO₂ = 2325 μ mol/l.

For this water the $[CO_3^-] = 75\mu M$ and $[HCO_3^2] = 2250 \mu M$

Calculate the new carbonate alkalinity, TCO₂, [HCO₃-] and [CO₃-] if the following occur:

a) 10 mg of foraminiferal shells are dissolved in 1 liter of the sample. (CaCO₃ molecular weight is 100 g/mole.

If 10 mg of CaCO₃ is dissolved 0.1mM or100 μmol CO3²⁻ is added;

So the new Alkalinity would be 2400 + 100x2 = 2600

The new $CO3^{2-}$ would be $75 + 100 = 175 \mu M$

The new $[HCO_3^-]$ would be Alk $-2[CO_3^{2^-}]$ thus 2250

New T_{CO2} will be 2325 + 100 = 2425 μ mol/l

b) 10 µmol of POM is completely respired in 1 liter of the seawater. You can use the simplifying assumption that one mole of dissolved carbon dioxide will react with one mole of carbonate ion to produce 2 moles of bicarbonate.

e.g.
$$CO_2 + CO_3^= + H_2O < ---> 2HCO_3^-$$

Respiration of 10 μ M of POM would release 10 μ mol CO₂ produced thus adding 10 μ moles to the T_{CO2}. There is no charge change so no change in alkalinity. T_{CO2} = 2325 + 10 = 2335

As can be seen from the above equation each mole of CO_2 produced will consume a mole of CO_3^{2-} so $75-10 = 65 \mu M$

 HCO_3^- on the other hand will increase be 2 moles for every mole of CO_2 produced thus $2250 + 20 = 2270 \mu M$

(8) How will fossil fuel CO₂ change the pH of the ocean? Write the equations

Scientists project that P_{CO2} will eventually double from its pre-industrial value of 280ppm to at least 600ppm it is now 370ppm.

Assume alkalinity stays constant at 2.300×10^{-3} eq 1^{-1} . DIC is now 1.95 (mmol/kg) Assume $K_H = 10^{-1.53}$ $K_1' = 10^{-6.0}$ and $K_2' = 10^{-9.1}$

- a. What will be the final pH for this model
- b. How valid is the assumption about const Alk?
- c. After 1000 years when the water mixed down deep how will this change CCD lysocline?

 $CO_2 + H_2O = HCO_3^{-} + H^{+}$ nothing else happening pH will decrease.

Using 370ppm today the $P_{CO2} = 10.91\mu M$. We can calculate from K_H and 600ppm that the P_{CO2} will be 17.7 μ M and the change in T_{CO2} (1950 + 6.79 =1956.8); from that get HCO_3^- (1613.6) and CO_3^{2-} (343.2); pH = 8.327 so it is lower by 0.012 pH units

The assumption is valid if no dissolution of CaCO₃ occurs

pH is a little lower and CO₃²⁻ is lower so there will be more dissolution and the CCD will become shallower.

(9) Going from the deep North Atlantic to the deep North Pacific, alkalinity increases in seawater from 2350 to 2475 μ Eq kg⁻¹ while total CO₂ increases from 2200 to 2375 μ M kg⁻¹. On average what is the relative contributions of CaCO₃ dissolution and organic matter respiration for producing these changes.

 Δ alk = 125, this is only due to CaCO₃ dissolution so 125/2 = 62.5 μ M CaCO₃ dissolved; Δ DIC = 175, this is a combination of CaCO₃ dissolution and OrgC oxidation. We can subtract the dissolution to evaluate the respiration: 175-62.5 = 112.5 μ M C form respiration; This gives a ratio of 1.8 orgC oxidized for 1 CaCO₃ dissolved.

(10) Using the elemental composition of marine particulate material as collected in sediment traps listed below (Broecker and Peng, 1982) (Relative to P = 1 mole) what changes in deep-ocean DIC and Alk do we predict if typical sediment trap particles respire and dissolve.

	P	N	С	Ca	Si
Soft Parts	1	15	105	0	0
Hard Parts	0	0	26	26	50
Composite	1	15	131	26	50

The Org.C : CaCO₃ in particulate matter is 105 : 26 (4 : 1)CaCO₃ + H⁺ = Ca²⁺ + HCO₃⁻ or $CO_2 + H_2O + CaCO_3 = Ca^{2+} + 2HCO_3$ ⁻ $O_2 + CH_2O = CO_2 + H_2O = H^+ + HCO_3$ ⁻

	1 mol CaCO ₃	4 mol Org. C.	Composite
ΔDIC	1	4	5
<mark>ΔCa</mark>	1	0	1
ΔAlk	<mark>2</mark>	0	2

DIC of the composite matter changes by 5 while Alk changes only be 2 so Δ DIC =2/5 Δ Alk, more change in DIC.

Alk_C – DIC = ([HCO₃⁻] + 2[CO₃²-]) – ([HCO₃⁻] + [CO₃²-] + [CO₂]) = [CO₃²-] – [CO₂] \cong [CO₃²-] Degradation of Org.C. and dissolution of CaCO₃ in the ratios in particulate matter cause a decrease in the carbonate ion concentration because Δ Alk – Δ DIC = Δ [CO₃²-] and alkalinity changes less than DIC.

 $a_{H+} = \{ [HCO_3^{-1}] / [CO_3^{2-1}] \} K_2' \text{ pH} = -\log a_{H+} [CO_3^{2-1}] \text{ down} => a_{H+} \text{ up} => \text{pH down} \}$